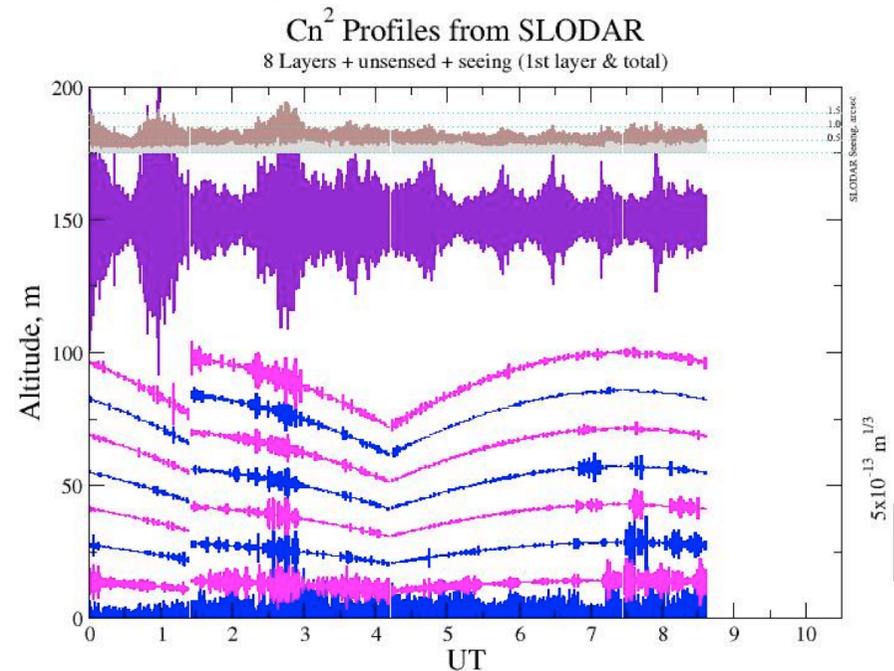


# Characterisation and Simulation of Atmospheric Seeing

Richard Wilson, Tim Butterley, Harry Shepherd  
CfAI, Durham University, UK



SLODAR Turbulence Profiler  
ESO Paranal Observatory



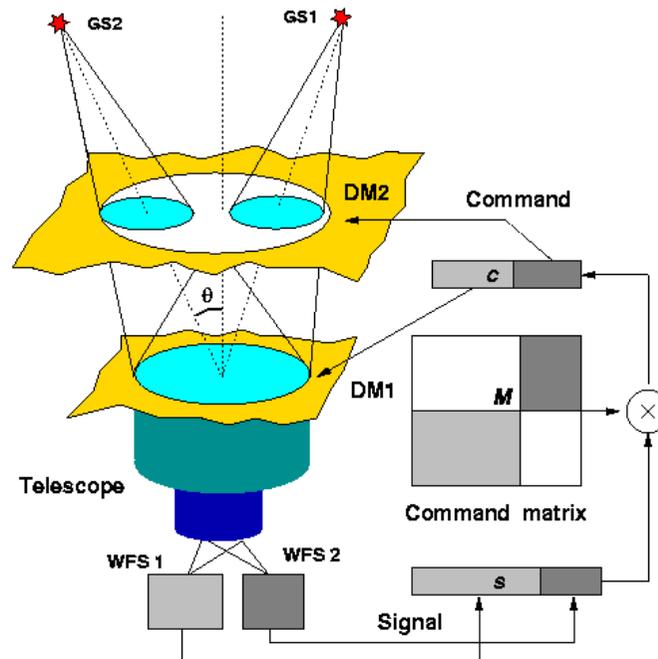
Turbulence profile sequence, Paranal



# Computer Simulation of Seeing

## The Usual Assumptions:

1. Thin layers, geometric propagation
2. Kolmogorov statistics for phase aberrations
3. Taylor approximation (fixed, translating pattern)
4. (Weak turbulence - neglect scintillation effects)



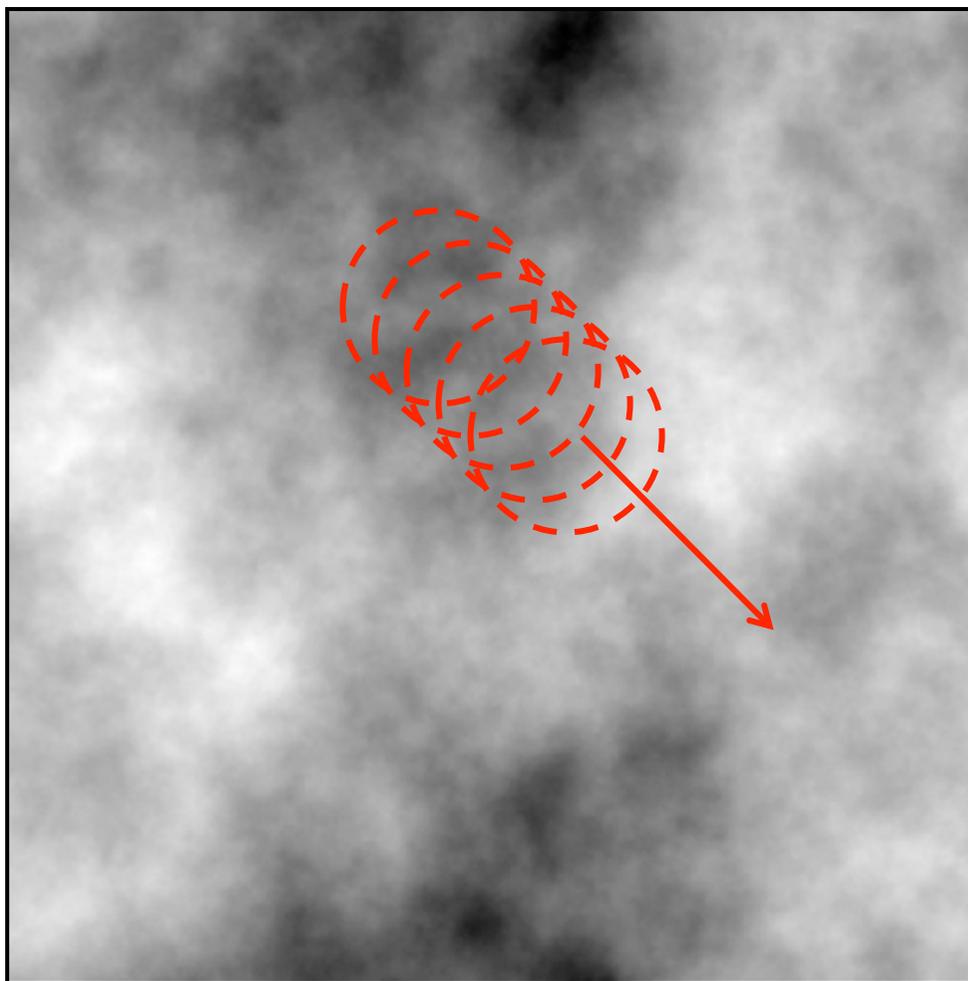
# Computer Simulation of Seeing

- **Creating a random turbulent ‘phase screen’**
- Power spectrum of phase aberrations after propagation through Kolmogorov turbulence:

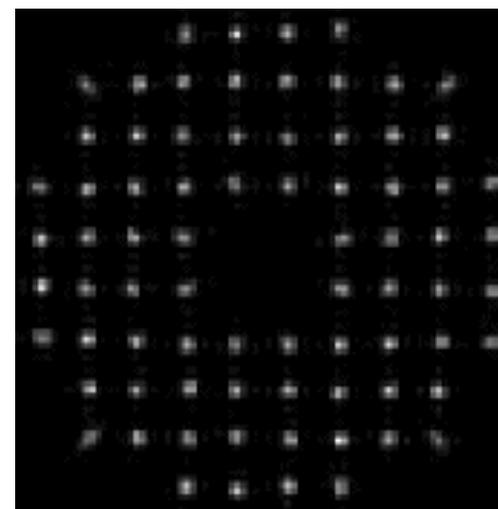
$$\Phi(\underline{k}) = 0.0229 r_0^{5/3} \underline{k}^{-11/3}$$

1. Create a 2d array of random numbers (normal distribution, mean = 0, variance = 1)
2. Multiply by  $\sqrt{\Phi(\underline{k})}$
3. Fourier transform to get an array of ‘phase’ fluctuations with the correct spatial structure function (almost).

# Computer Simulation of Seeing



Kolmogorov 'phase' screen



Simulated Wavefront-sensor pattern

# Computer Simulation of Seeing

- **Creating a speckle image:**

The instantaneous PSF is the squared Fourier transform of the (complex aperture) function  $T(\rho)$  :

$$I(x) = \mathbf{F}^2[T(\rho)]$$

1. Create a complex array to contain  $T(\rho) = Ae^{i\phi(\rho)}$
2. Set phase of  $T(\rho)$  equal to values from Kolmogorov phase screen
3. Amplitude of  $T$  is constant within a aperture (zero outside)
4.  $\text{FFT}^2$  gives speckle image
5. Now move the aperture across the phase screen to simulate wind-blown motion of the turbulence...

# AO Numerical Simulations

- Monte Carlo simulations of AO typically make use of this basic method, but have become highly complex:
  - Multiple WFS and science ‘light paths’, e.g. MOAO
  - NGS and LGS propagation
  - WFS noise, DM hysteresis, etc...
  - 8m/ELT scale AO -> Parallel programming

# What is the Atmosphere Really Like ?

1. Thin Layers – incl. surface layer
  - Is isoplanatism modeled correctly ?
2. Kolmogorov statistics (spatial structure)
  - Is DM fitting error modeled correctly ?
3. Taylor approximation ?
  - Are temporal effects modeled correctly ?

# Turbulence Characterisation

## (Ideal) Requirements for a Turbulence Profiler:

- Measure  $C_n^2(h)$ , +  $V_w(h)$  if possible
- Measure all altitudes with high resolution
- Well calibrated (turbulence strength, altitude)
- Good time resolution (~1 minute)
- Real-time data
- Automated / robotic
- Portable (re-locatable) for site testing
- Cheap

# Turbulence Profiling: Methods

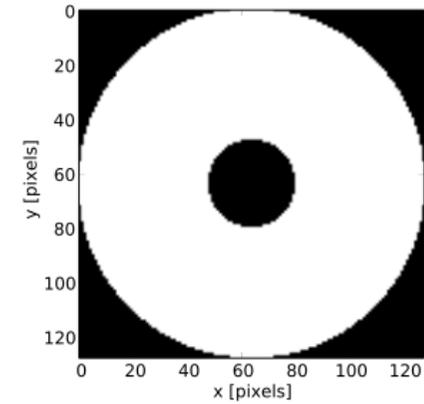


Micro-Thermal Balloon Probe  
Measures  $C^2_T(h)$

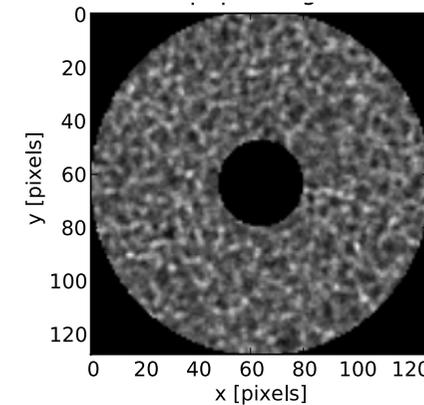


Acoustic Ranging (SODAR)

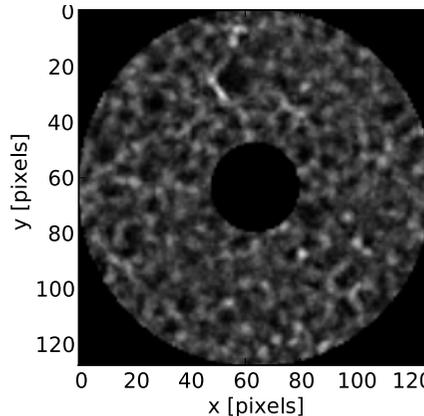
# MASS: Scintillation Spatial Structure



0km



5km



10km

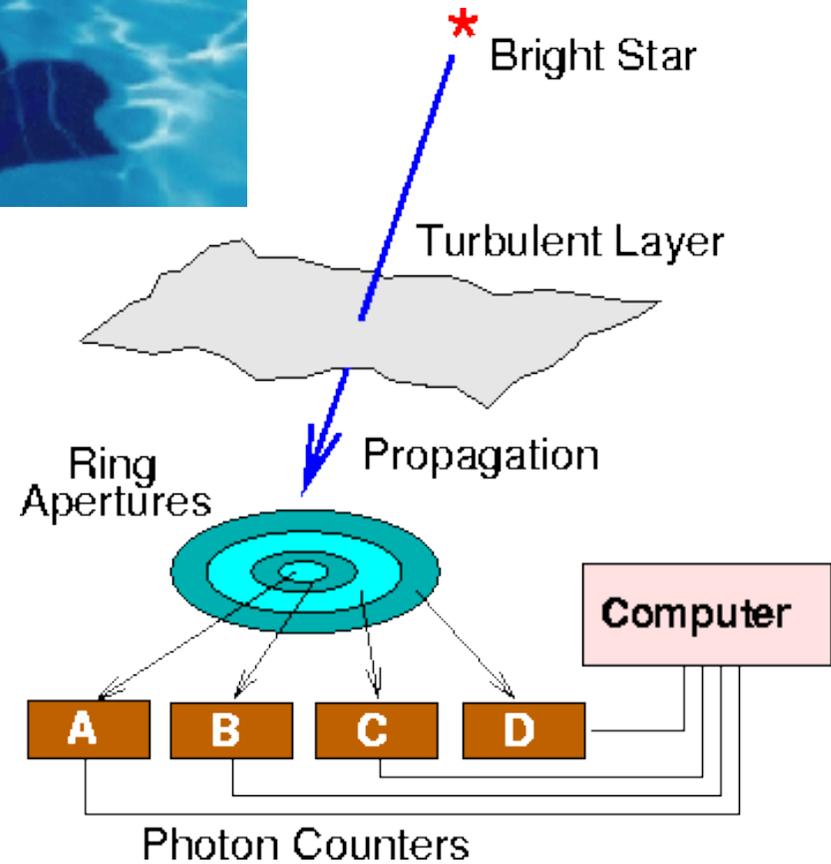


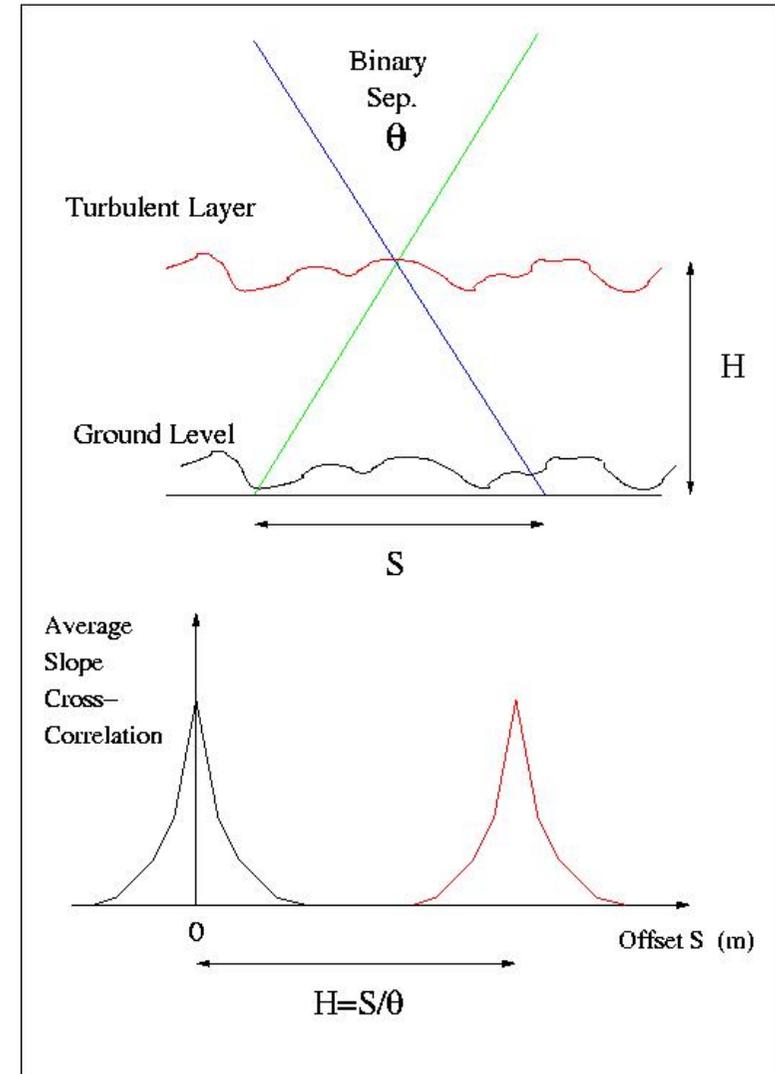
Image: <http://www.ctio.noao.edu/~atokovin/profiler/index.html>

# Optical Crossed-Beams Methods

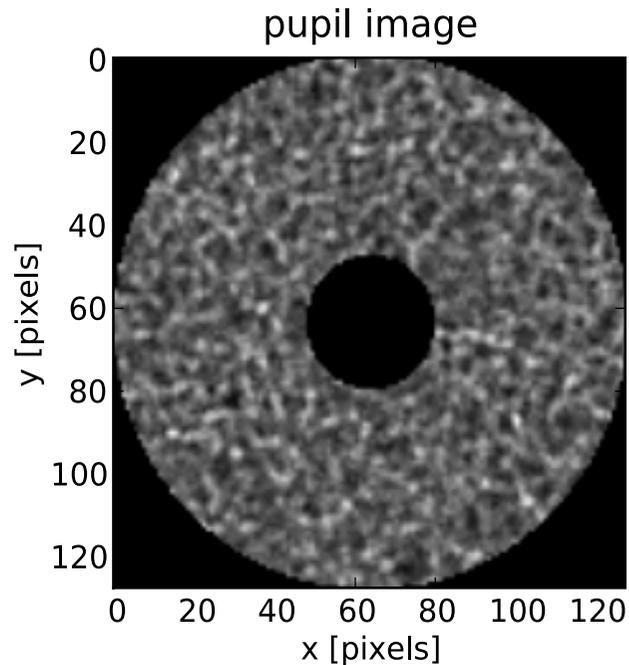
## Recipe:

1. Observe a double star
2. Measure something at the ground... scintillation pattern, WFS centroids
3. Recover  $C_n^2(h)$  from the **time-averaged cross-covariance** of the data

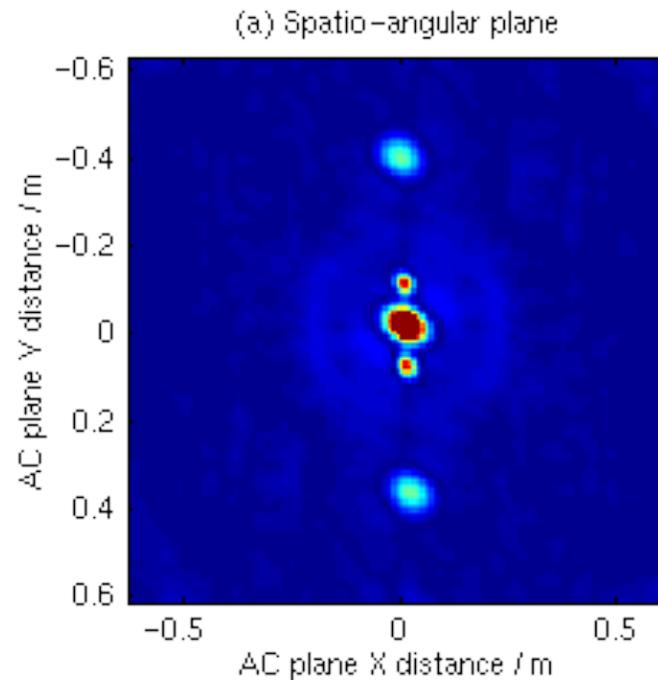
e.g. SCIDAR, LOLAS, SLODAR, Co-SLIDAR ....



# SCIDAR (SCIntillation Detection And Ranging)



Double Star  
Intensity Pattern

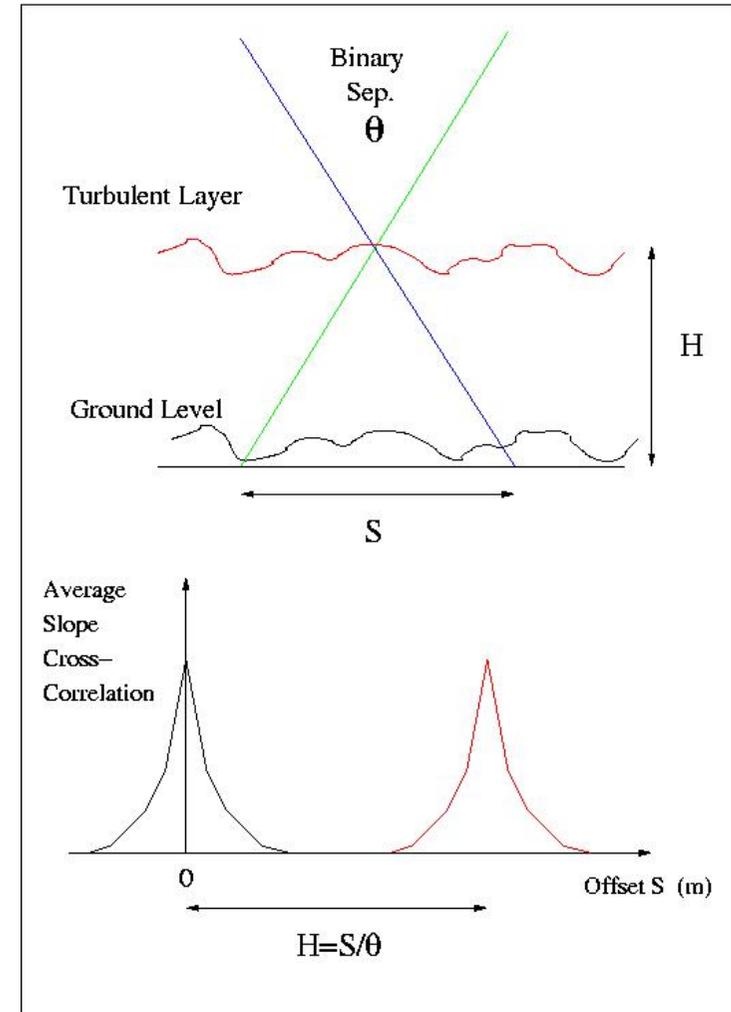
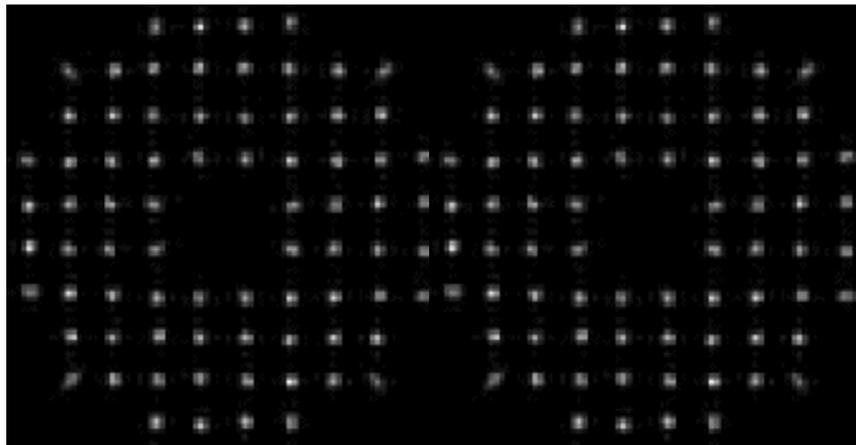


Intensity Auto-Correlation  
(Time-Averaged)

Altitude resolution determined by double star separation and Fresnel zone size (roughly)

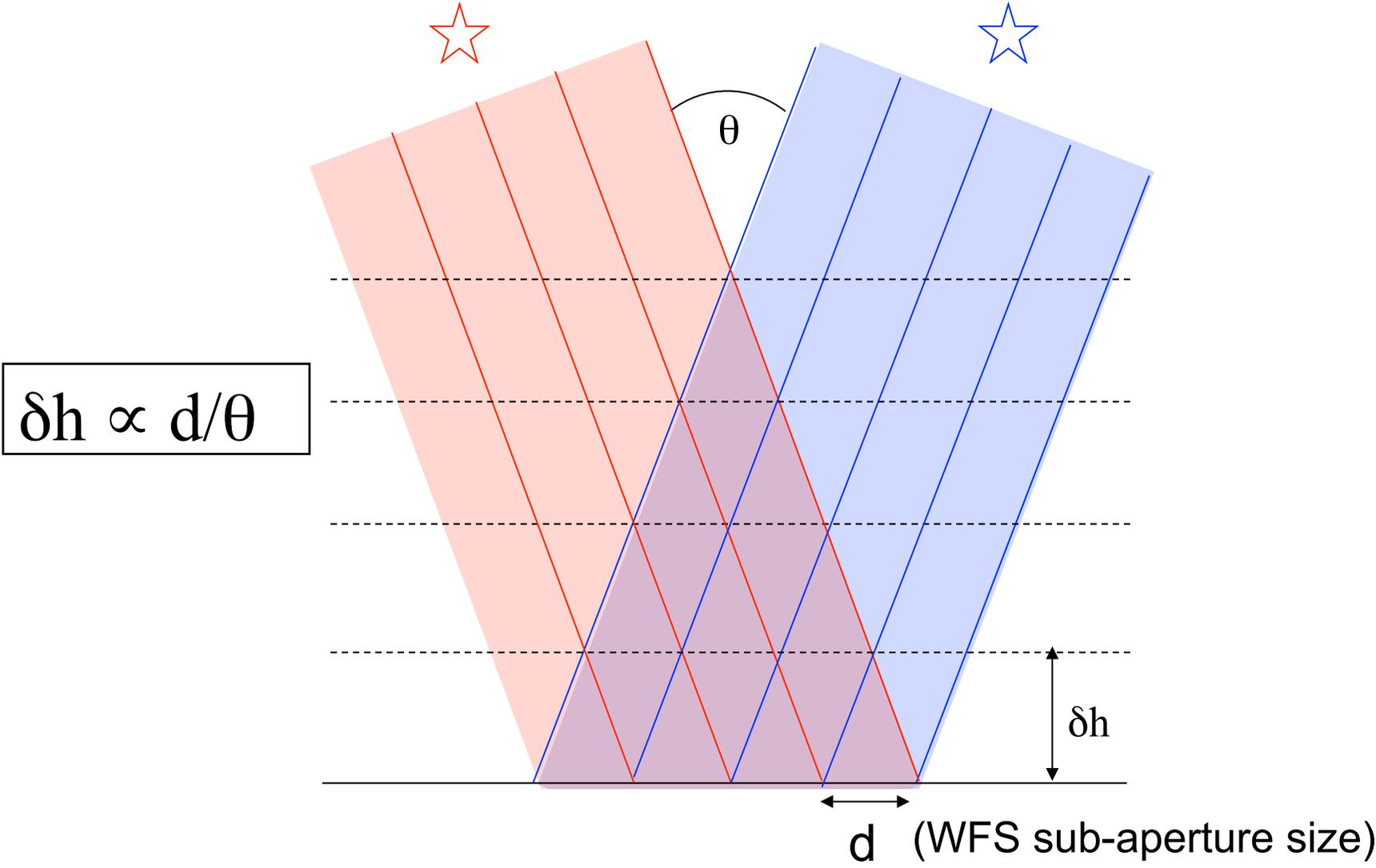
# SLODAR (SLOpe Detection And Ranging)

- Shack-Hartmann wavefront sensor
- Recover  $C_n^2(h)$  from the time-averaged cross-covariance of the WFS data

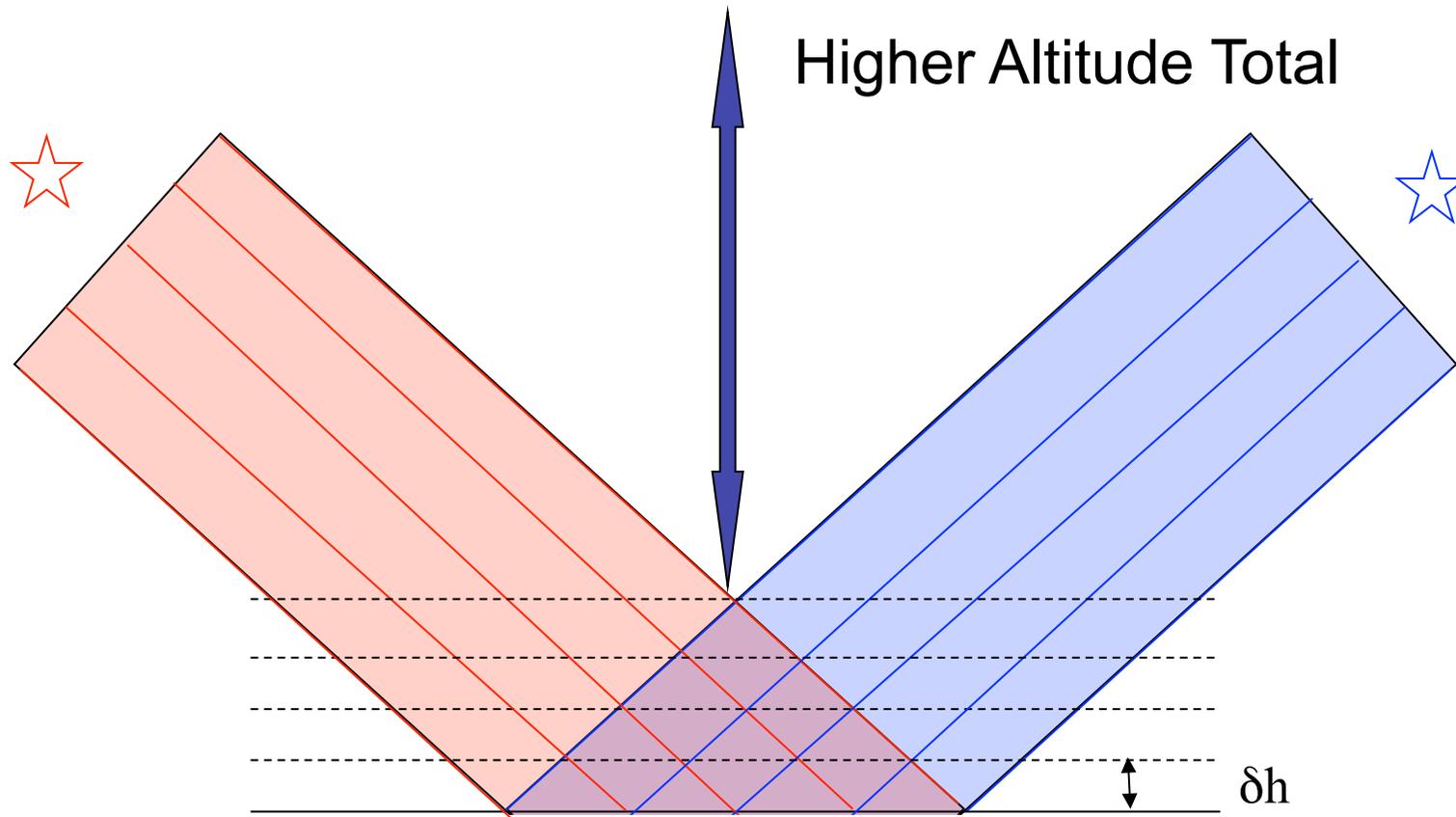


Altitude resolution determined by double star separation and Sub-apertures size (roughly)

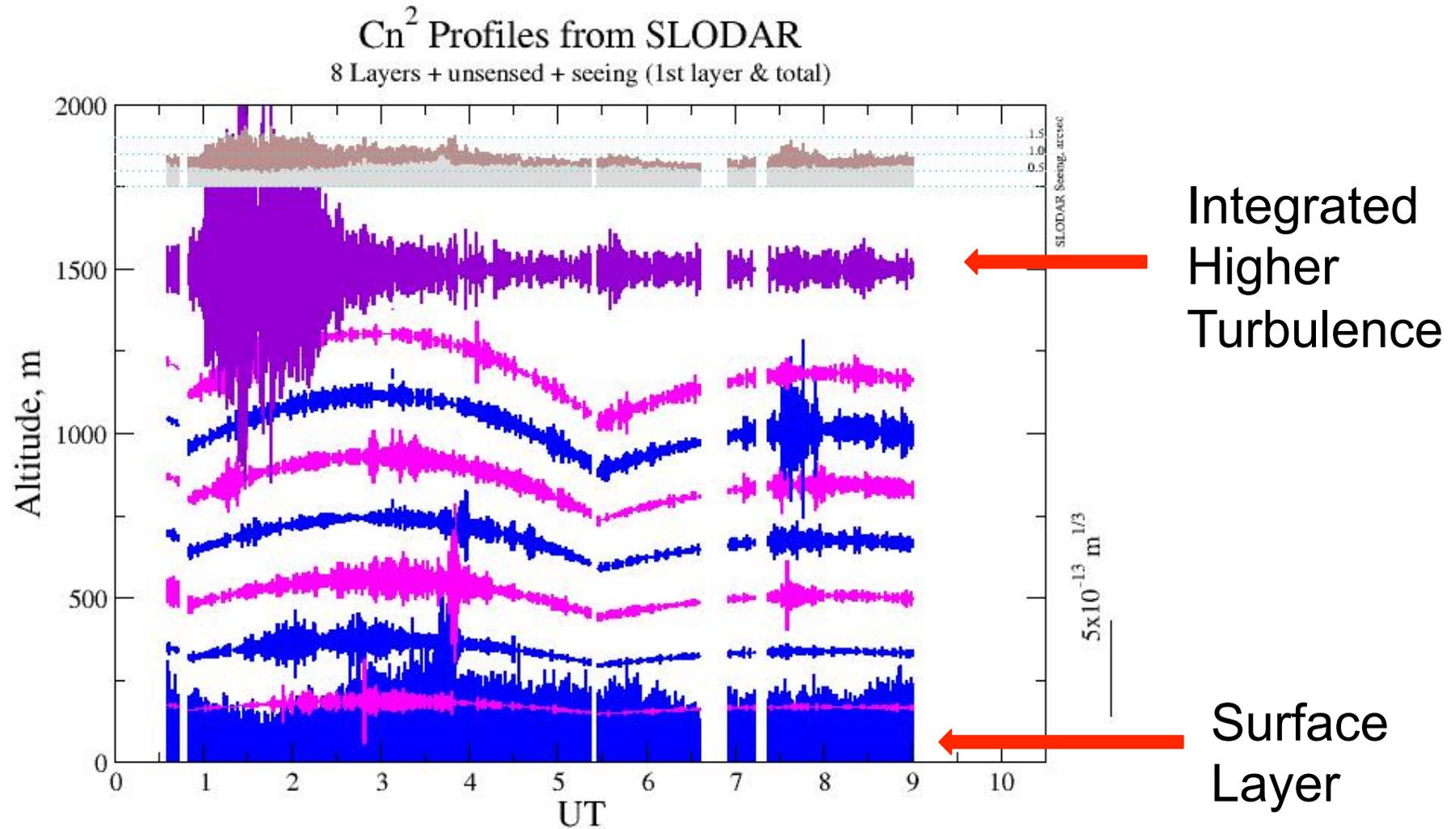
# SLODAR Altitude Sampling



# SLODAR Altitude Sampling: Ground-Layer profiling



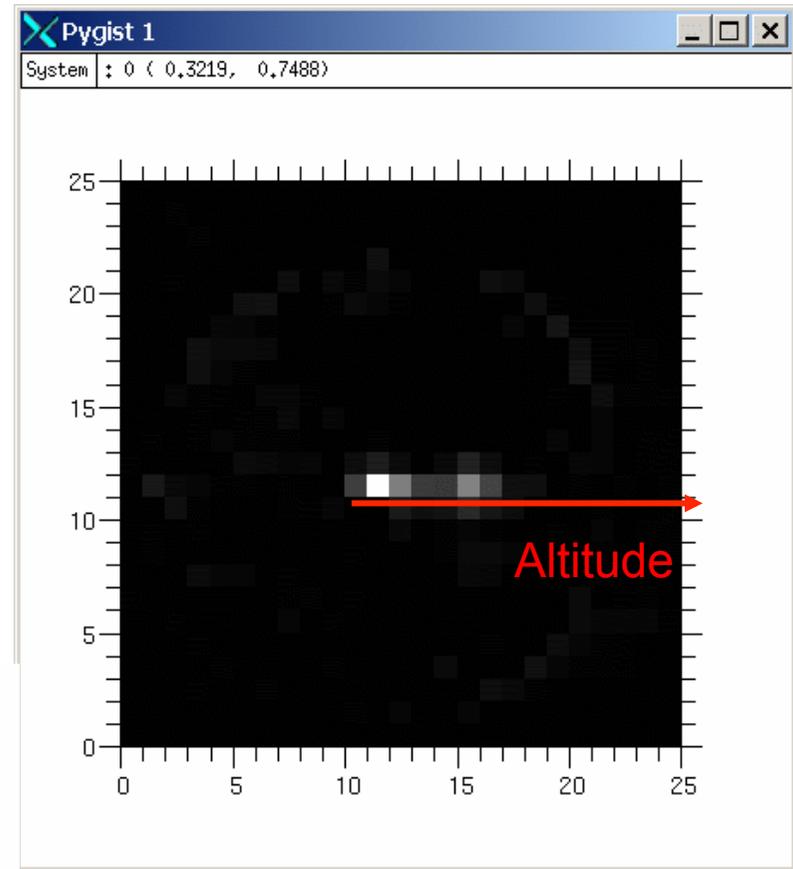
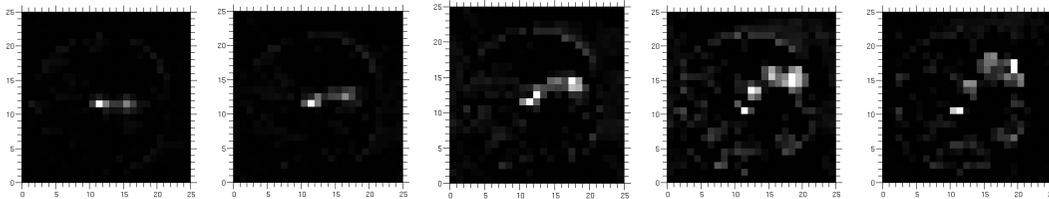
# SLODAR Turbulence Profile Sequence (Paranal)



Altitude resolution varies with zenith angle

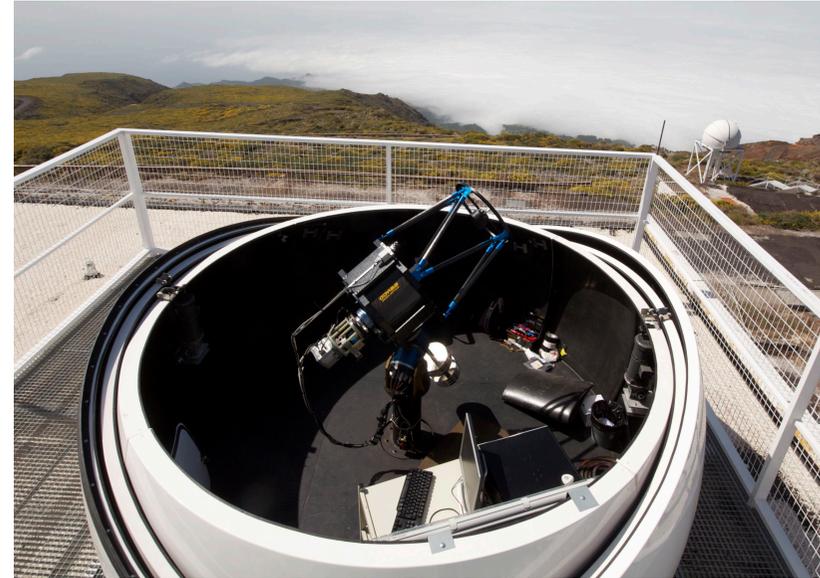
# SLODAR: Turbulent Layer Velocities

- Movie shows the spatial cross-covariance with increasing time offset.
- Motion of peaks => layer velocity (with altitude).



# SLODAR at La Palma

- Support CANARY MOAO demonstrator at WHT
- 0.5m telescope
- ‘Whole’ atmosphere profiling (up to ~10km)
- Fully automated



# SLODAR Site Monitor, ESO Paranal, Chile



# Surface-Layer SLODAR, Paranal

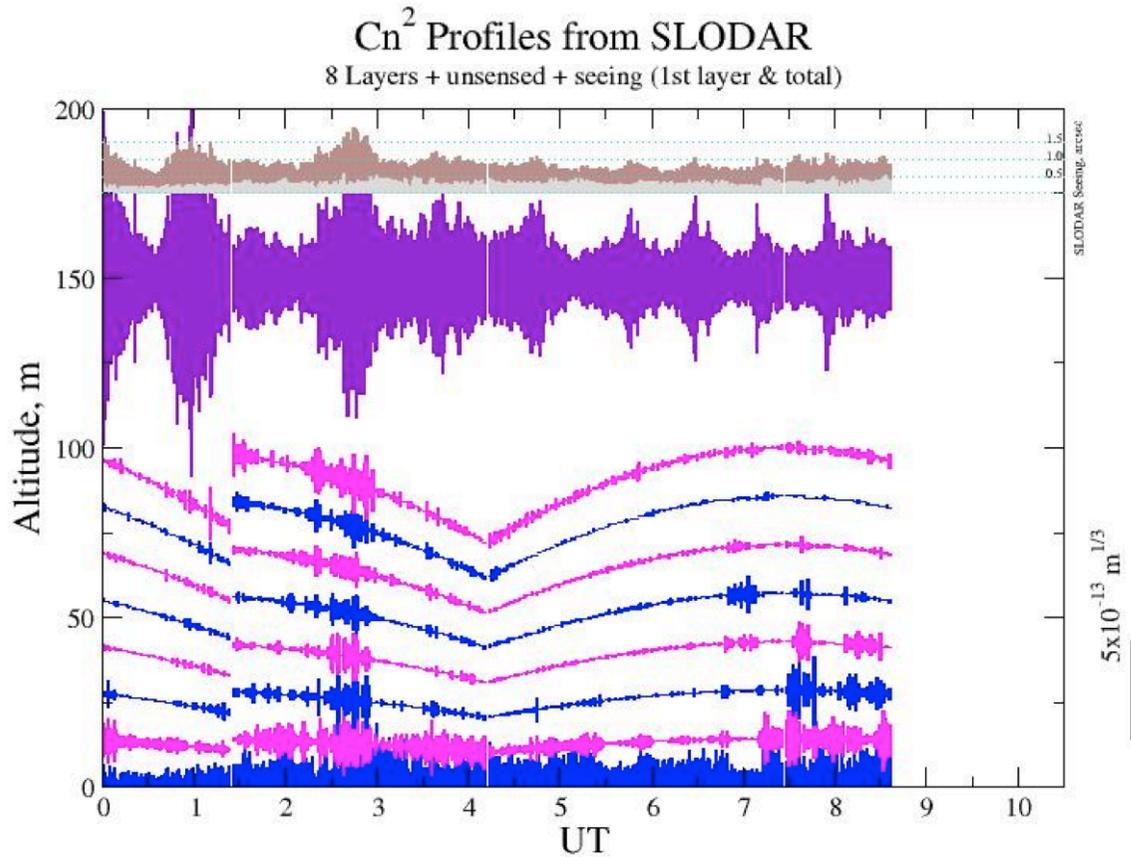


Photo: Tim Butterley

# How Accurate are the Usual Assumptions ?

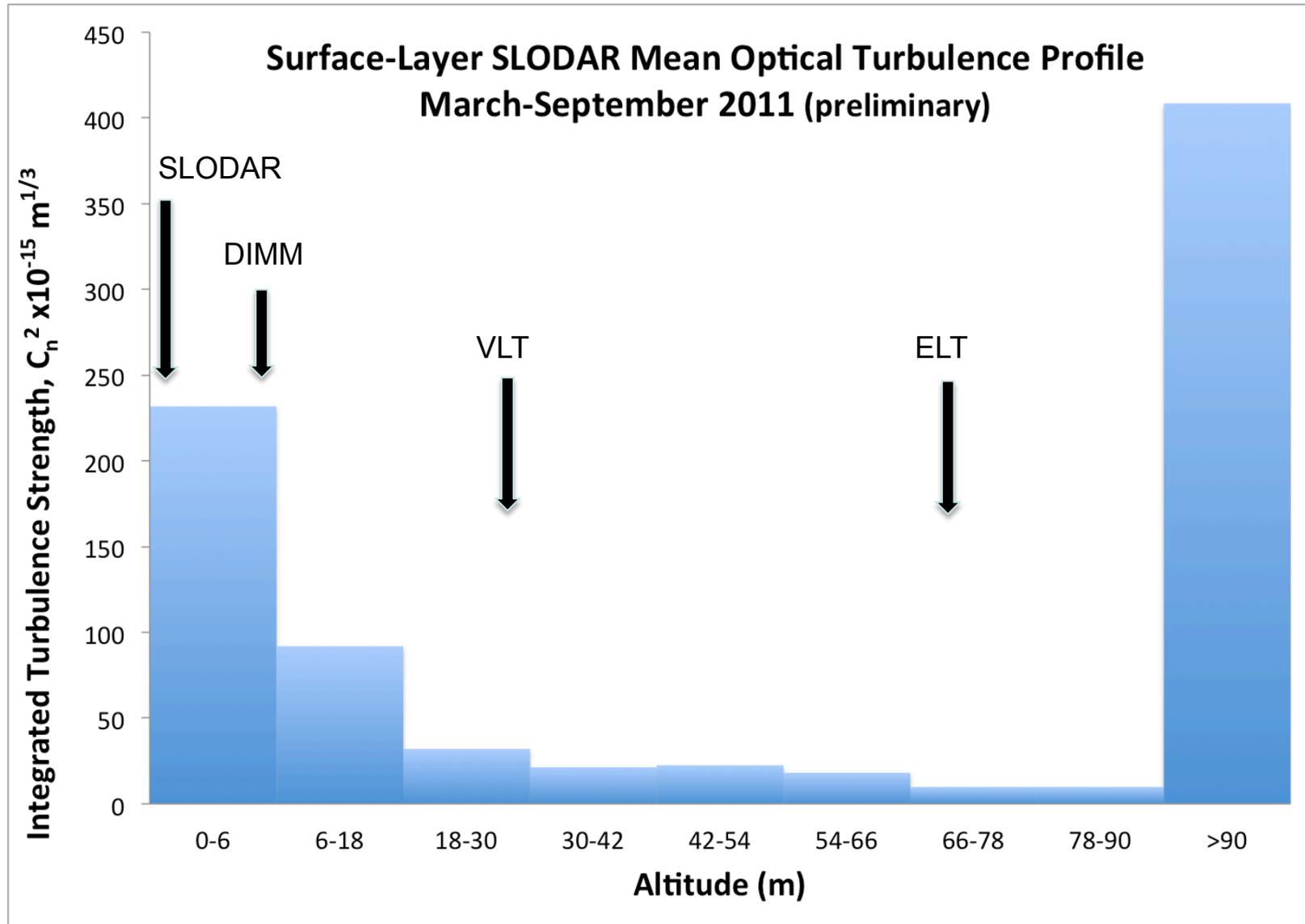
## 1. Thin Layers ?

Surface layer

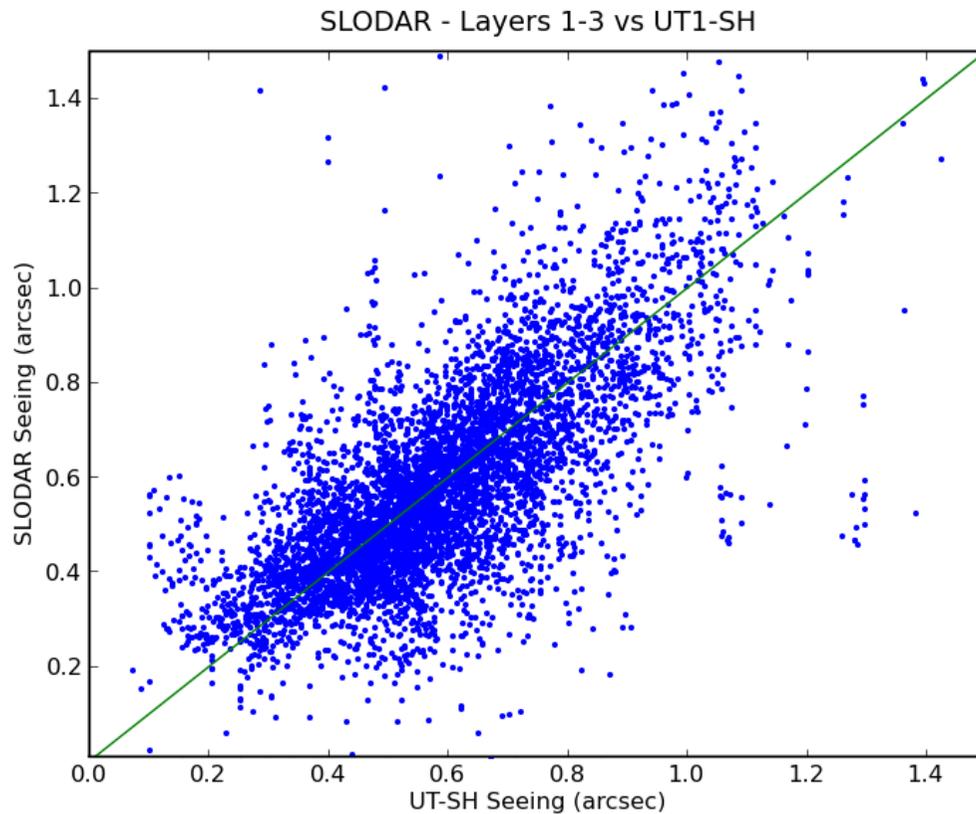


Turbulence profile sequence, Paranal

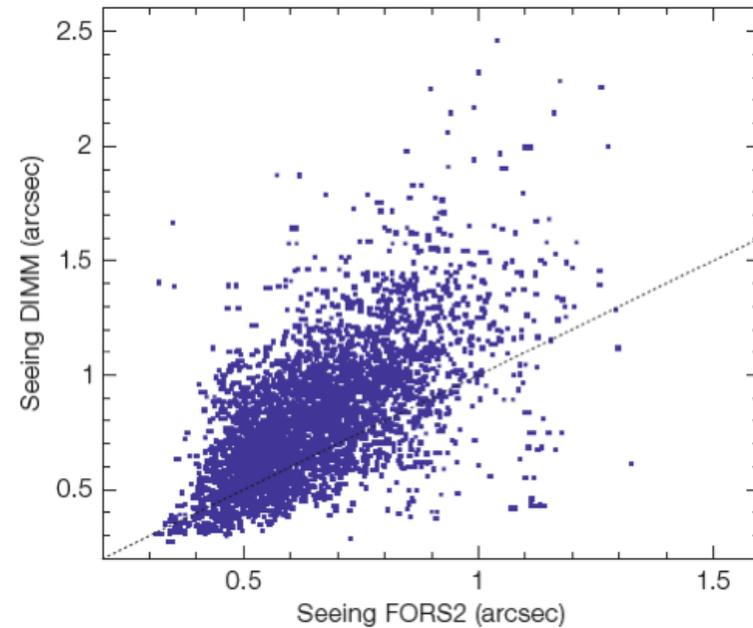
# Paranal: Surface Layer Turbulence Profile



# Understanding VLT (UT) Seeing



SLODAR total above 30m  
versus UT shack-Hartmann seeing.



Paranal DIMM (at 5m)  
versus UT seeing

# Ground-Layer Turbulence: Does GLAO work ?

Thin surface layer –  
no need for GLAO

Thicker surface layer –  
GLAO is very effective and  
has a huge field of view

How much turbulence does  
the telescope create ?

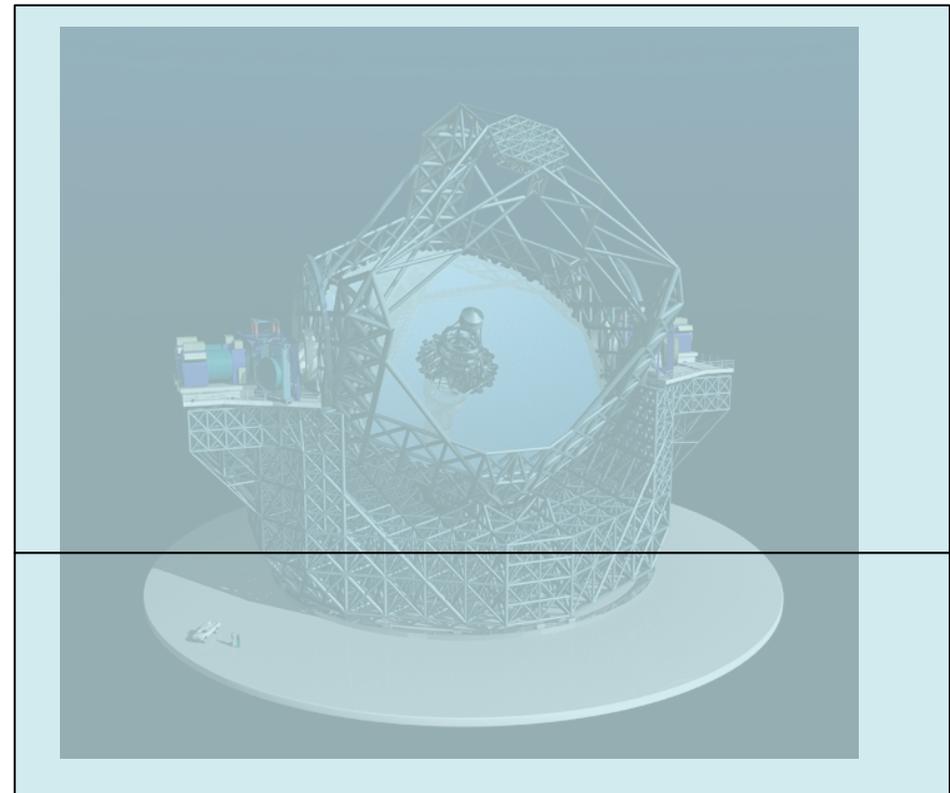
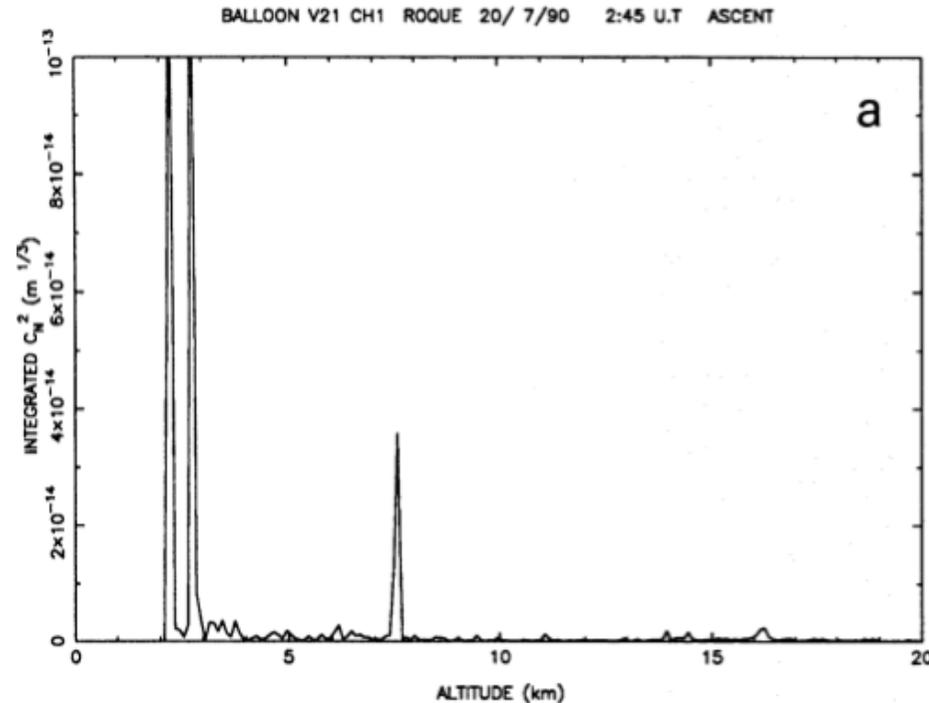


Image: [www.eso.org](http://www.eso.org)

# How Accurate are the Usual Assumptions ?

## 1. Thin Layers ? High layers

Very little high resolution profile data available (e.g. balloon soundings)



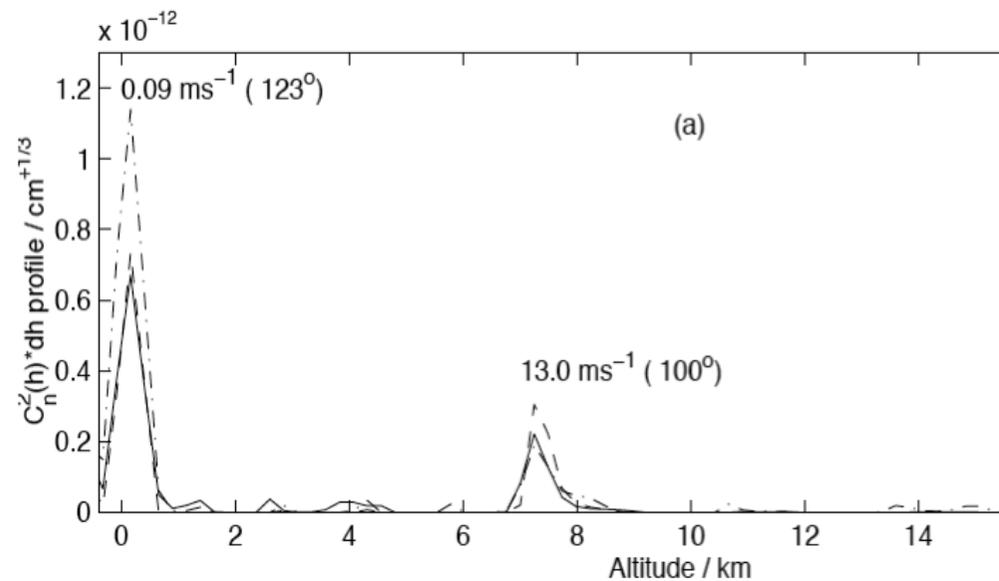
Balloon sounding turbulence profile,  
La Palma, Vertical resolution ~50m

Vernini et al, AA 204, 311, 1994

# How Accurate are the Usual Assumptions ?

## 1. Thin Layers ? High layers

Lots of low resolution profile data available (e.g. SCIDAR)



SCIDAR turbulence profile, La Palma  
Vertical resolution  $\sim 500\text{m}$

Kluckers et al, A&AS 130, 141, 1998

# How Accurate are the Usual Assumptions ?

## 1. Thin Layers ? High layers

- Implications: depends on the field of view...
- Individual layers are typically unresolved in low resolution data ( $\delta h \sim 500\text{m}$ ). No implications for modeling AO with small field of view  $< 1\text{arcmin}$  ?
- Higher resolution data,  $\delta h < 100\text{m}$ , may be critical for larger FOV, e.g. MOAO / EAGLE (5 arcmin)

# How Accurate are the Usual Assumptions ?

## 2. Kolmogorov statistics ?

Kolmogorov spatial spectrum:

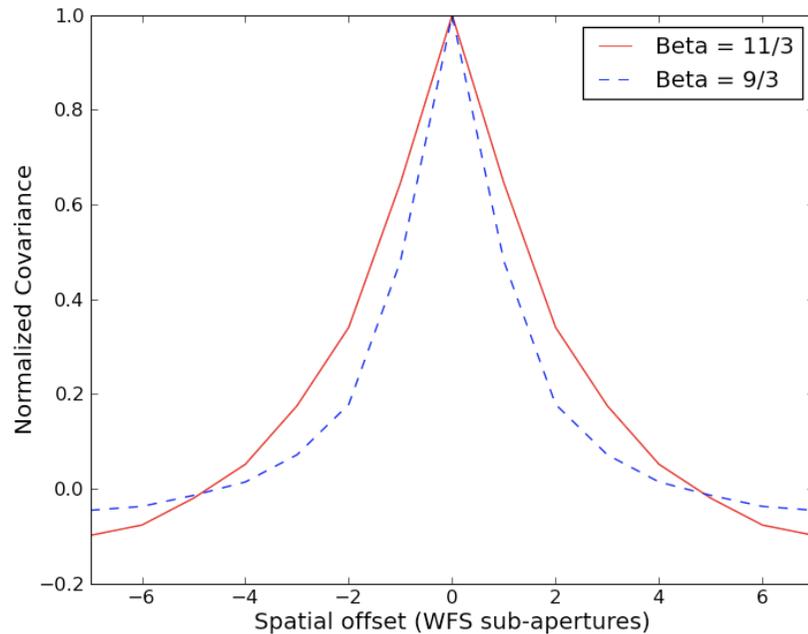
$$\Phi(\underline{k}) = 0.0229 r_0^{5/3} \underline{k}^{-\beta}, \quad \beta = 11/3$$

Surface Layer / 'local turbulence':

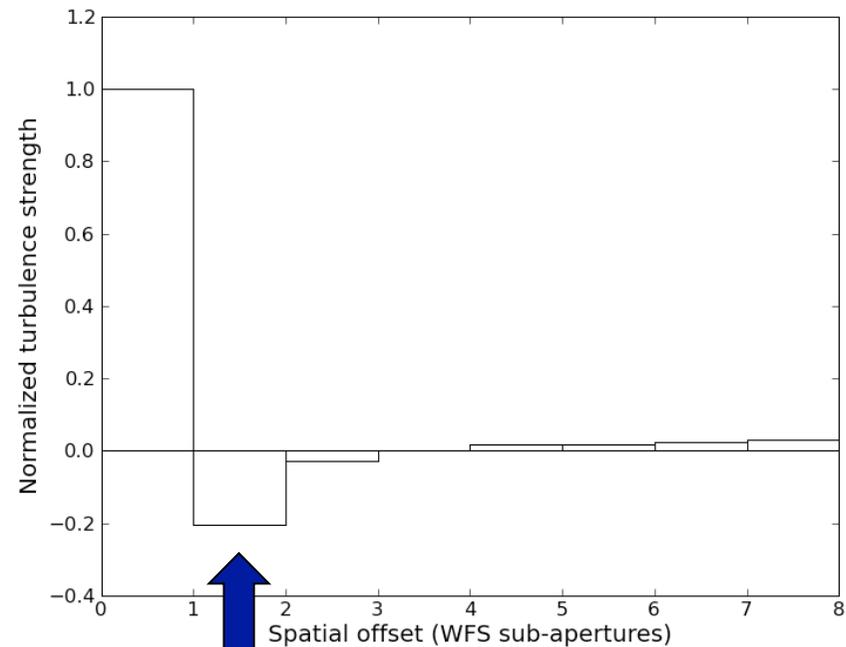
Often apparently measure  $\beta < 11/3$ , mainly in low wind speeds  
(Can also model as a small 'outer scale of turbulence')

**BUT: observed  $\beta < 11/3$  results because, in light winds, larger spatial scales are not properly sampled (unless our sampling time is very long...)**

# Non-Kolmogorov Response



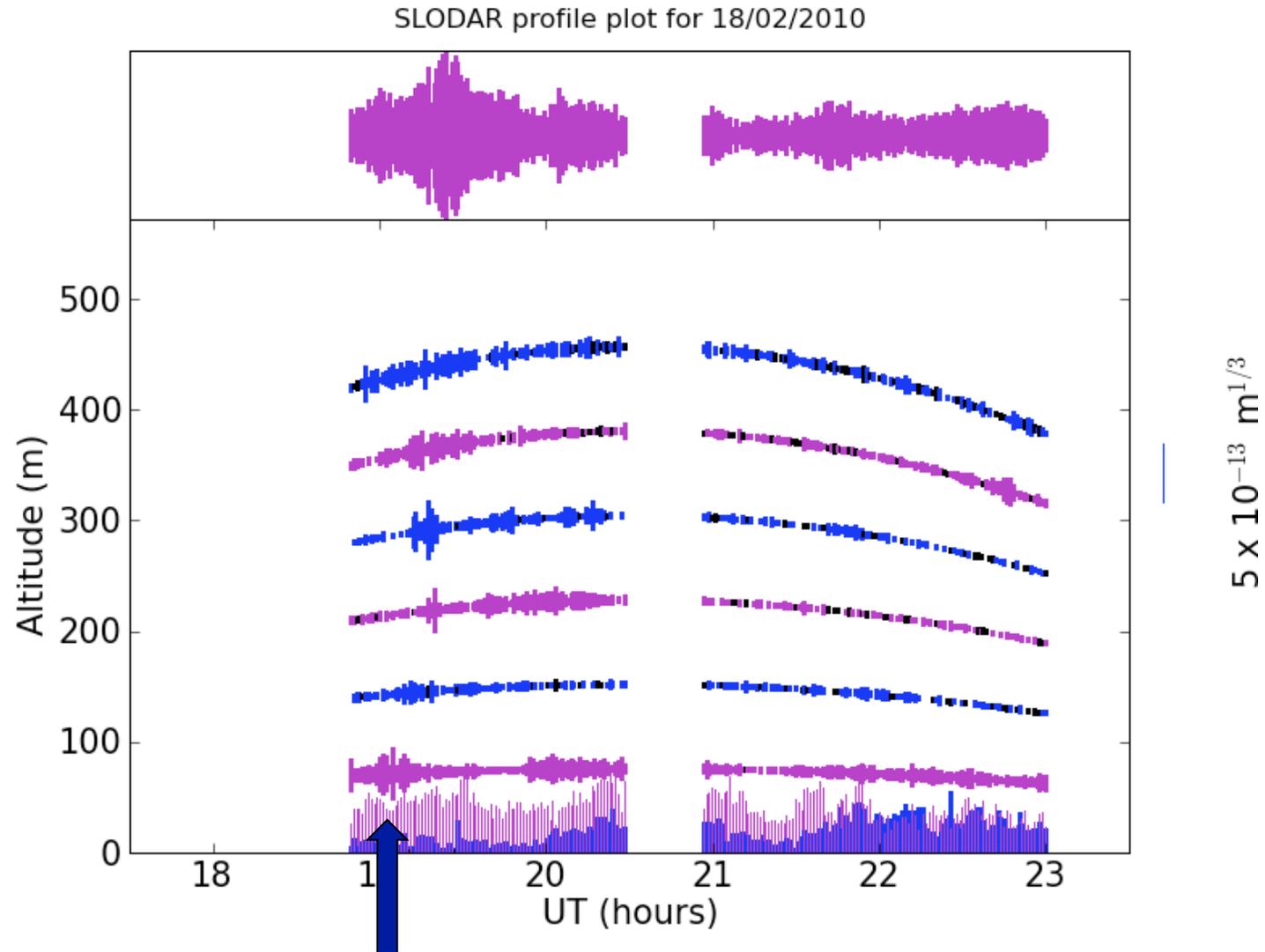
SLODAR response functions



Negative 'side-lobe' in restored turbulence profile

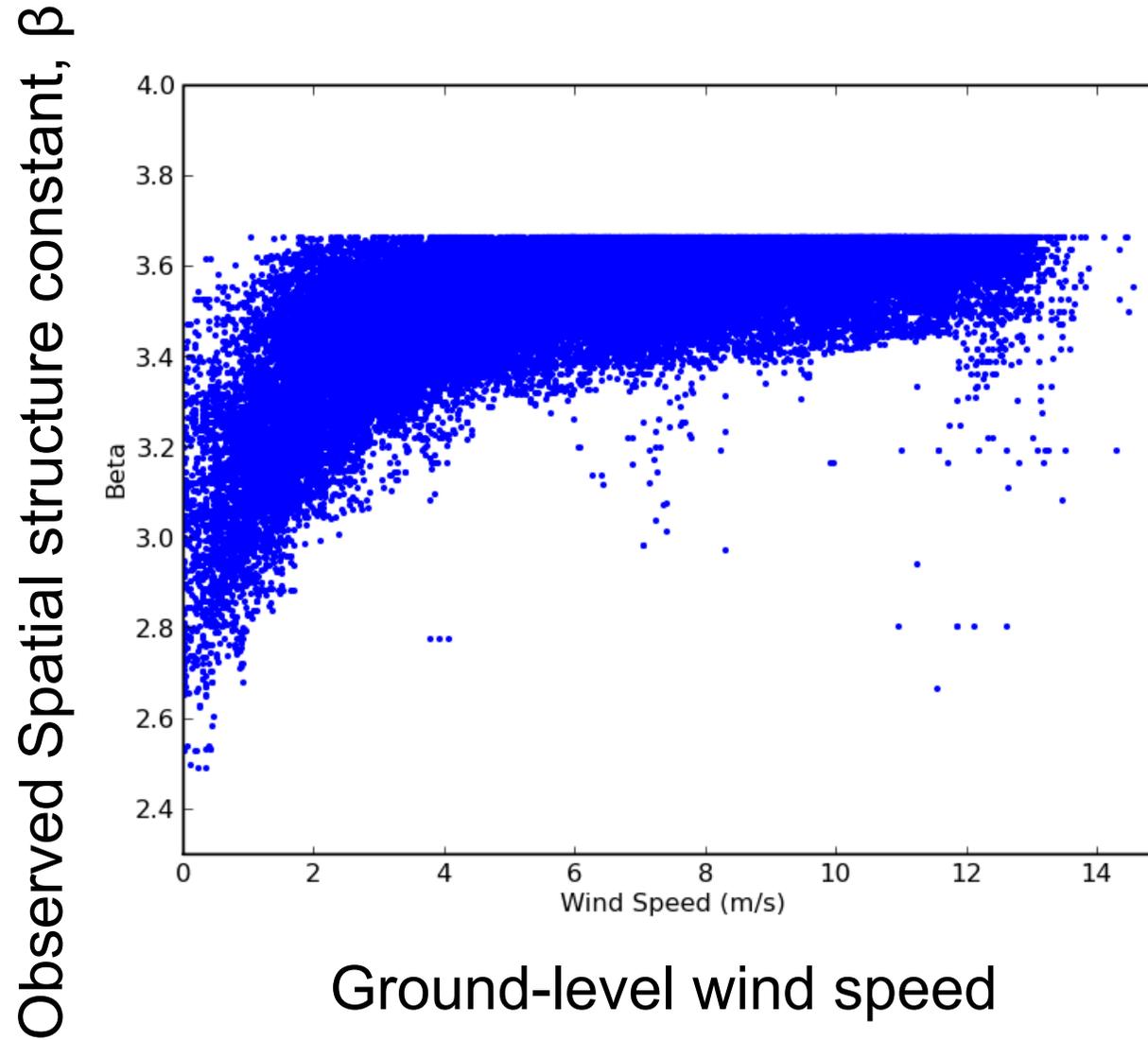
**Beware of Enforcing Positivity !**

# Non-Kolmogorov Response



'Fix': Include additional  $\beta < 11/3$  term at the ground

# Non-Kolmogorov Response



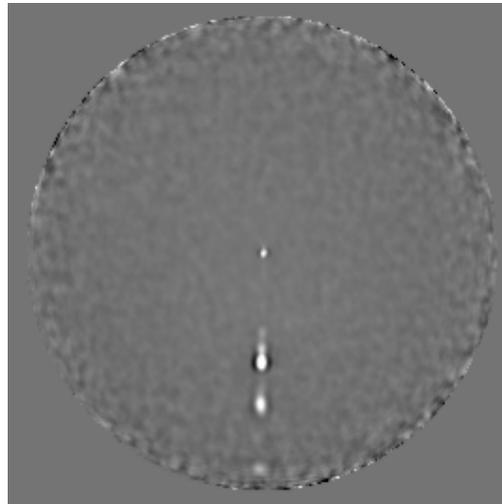
# How Accurate are the Usual Assumptions ?

## Taylor approximation

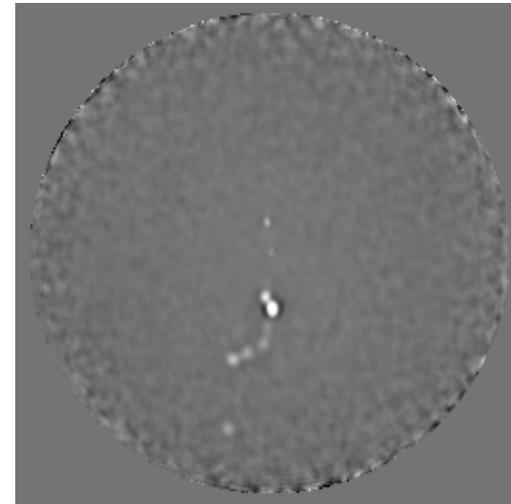
- Turbulence results from wind-shear => we should expect velocity dispersion
- High altitude- and time- resolution data needed...

High-resolution SCIDAR data, Nordic Optical Telescope, La Palma

Images: Harry Shepherd



$\Delta t = 0\text{ms}$



$\Delta t = 50\text{ms}$

# How Accurate are the Usual Assumptions ?

1. Thin Layers – incl. surface layer
  - ✗ • Surface-Layer structure is critical
  - ✓ (?) • High layers: may break down for very large field of view ?
2. Kolmogorov statistics (spatial structure)
  - ✓ • Probably OK, but:
  - Need to model layer velocities correctly (e.g. range of surface wind-speeds)
3. Taylor approximation ?
  - ✓ (?) • More high-resolution profile data needed

# Key Issues for Profiling

- Surface-Layer Characterization
- Correct velocity structure
  - Effect of low winds (esp. surface layer)
  - Wind-shear
- Accurate Altitudes of (High) Layers:
  - Requirements for MCAO with ELT:  $<100\text{m}$  ?
  - How to do this (for the whole profile) ?